

AN OVERVIEW OF THE MINERALOGY OF THE
NEZILOVO, MACEDONIA DISTRICT AS COMPARED TO THE
FRANKLIN, NEW JERSEY DEPOSITS

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INTRODUCTION

The Franklin-Sterling Hill district in New Jersey has long been noted for its unique mineralogy and minable concentrations of zinc, iron, and manganese. In recent years the Nezilovo district in southern Yugoslavia has come under some scrutiny as possibly being a deposit of similar origin and composition. Little work seems to have been done in the Nezilovo area. From optical and X-ray examination of several rock specimens collected from this locale, the basic mineralogy and possible relationship to Franklin will be discussed.

FRANKLIN SUMMARY

The Franklin-Sterling Hill area lies about 50 km west of New York City just on the western boundary of the New Jersey Highlands and is bounded by $74^{\circ}30'$ and $74^{\circ}45'$ longitude and 41° and $41^{\circ}15'$ north latitude lines.

The deposits of this area first came to attention in 1640 shortly after the establishment of New Amsterdam. Little was done in the area until 1848 when the Sussex Zinc and Copper Mining and Manufacturing Company was created with the purpose of exploiting these deposits. In 1897 after much litigation over conflicting titles, this company and several others united to form the New Jersey Zinc Company

which still actively mines this region.

In 1954 the Franklin mine was completely worked out. It is estimated that 22-23 million tons of ore were recovered from the Franklin mine over its years of production. This ore averaged 19.6 percent Zn, 8.7 percent Mn, and 17 percent Fe. (Fronde1, 1974, p. 157.) The nearby Sterling Hill mine began production shortly after 1912 and is still a source of high grade zinc ore. In 1973 alone, New Jersey records having produced 33,027 short tons of zinc with Sterling Hill being undoubtedly a major supplier. (Minerals Yearbook, p. 1329.)

The most widely accepted theory of origin of the Franklin deposits is that of Baum and Fronde1 (1974). Some of the original sediments were terrigenously derived and deposited in a marine environment. These are now analogous in many respects to metamorphosed sedimentary manganese deposits found elsewhere. Evidence shows that other constituents were probably chemical sediments formed by submarine hot springs. All of these sediments were deposited contemporaneously with a layer of limestone. During a period of high-grade regional metamorphism in the PreCambrian these layers were recrystallized, folded and faulted. During the late PreCambrian the region was subjected to erosion and was then later overlain by a series of Paleozoic sediments; the Cambrian Hardyston conglomerate and the Ordovician Kittatinny limestone. In the Paleozoic, the area was faulted and some minor folding occurred. This was followed

by the intrusion of low-silica igneous dikes in the Triassic.
(See Table 1.)

Age	Formation	Remarks	Thickness
Cambro-Ordovician	Kittatinny	Mostly dolomitic limestone, in part shaly or sandy	2,500-3,000 ft?
Lower Cambrian	Hardyston	Conglomerate and quartzite	30 ft ±
Precambrian	Pochuck Gneiss Series	Interlayered hornblende gneiss, microcline gneiss, locally with garnet or graphite, also local quartzite	Over 2,000 ft
	Wildcat Marble	Coarsely crystalline marble	300 ft
	Cork Hill Gneiss	Similar to Pochuck Series	800-1,000 ft
	Franklin Marble	Coarsely crystalline marble, locally dolomitic	1,100-1,500 ft
	Median Gneiss	Chiefly biotite gneiss	50-300 ft
	Franklin Marble	As above	Unknown, over 500 ft?
	Hamburg Mountain Gneiss	Similar to Pochuck Series, intruded by Byram hornblende granite	Over 2,000 ft?

Data from Hague et al. (1956), Baker and Buddington (1970), and Spencer et al. (1908). (Friedel and Brown, p. 140)

The Franklin and Sterling ore bodies both dip about 55° to the southeast.

At Franklin the ore body assumes the shape of a synclinal fold with unequal axes. The fold strikes northwest and plunges 25° southeast. The west limb of the syncline has the longest exposure, 8 km. It varies in thickness from 4.5 m to 27 m, generally thickening toward the keel. The eastern limb has the shorter exposure of 183 m giving the outcrop a hook-shaped appearance.

The ores lie in the Franklin limestone within a few meters of the eastward-dipping contact between the Franklin limestone and the adjacent gneisses on the west. (See Figure 1.)

At Sterling the east limb of the fold is longest, parallels the west, and is overturned. Again contained in

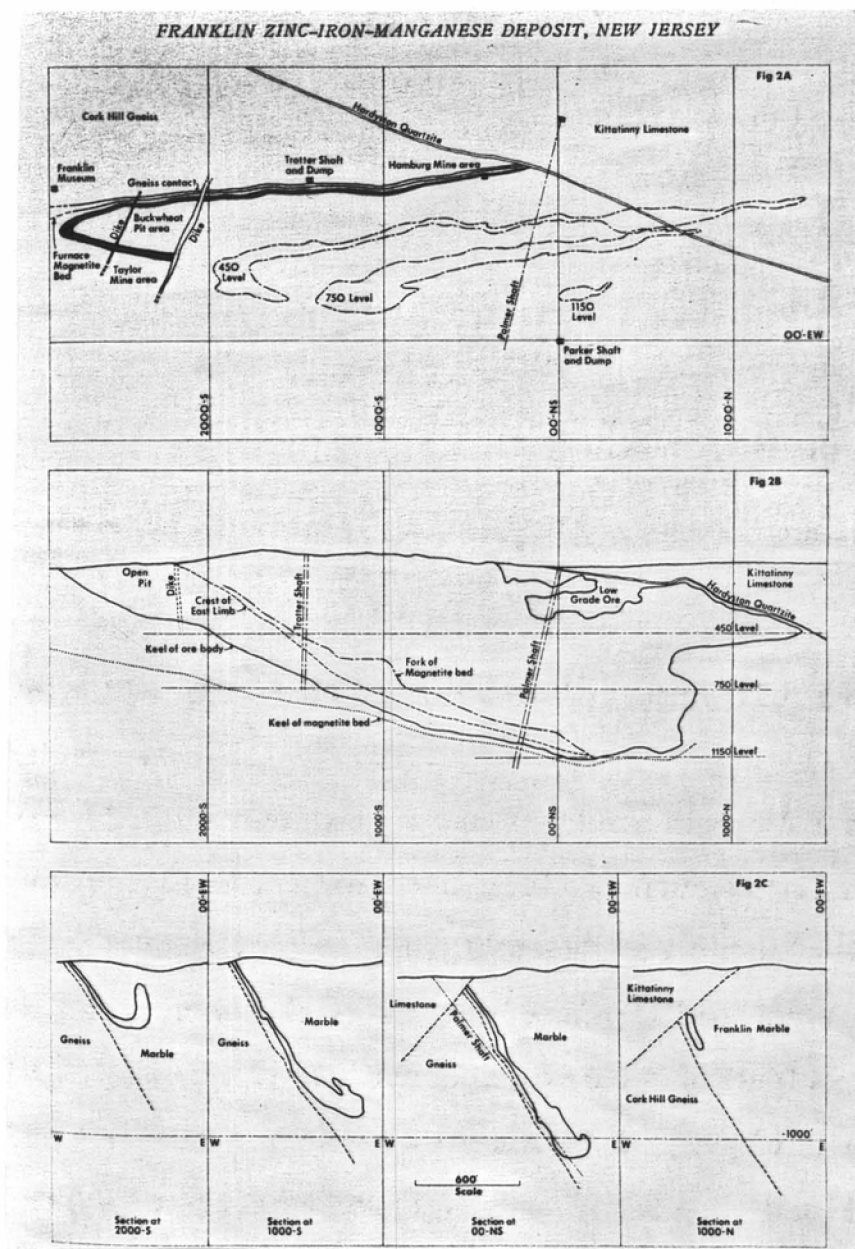


FIGURE 1. Plan, longitudinal elevation, and vertical cross sections of the Franklin ore body. Map coordinates north-south and east-west; zero coordinate at Parker Shaft. Scale in feet. Figure 2B is projected on an arbitrary north-south vertical plane. Adapted from maps by A. W. Pinger (1950) and C. H. Stockwell, New Jersey Zinc Company, 1951. (Fronzel and Baum, p. 161)

the Franklin limestone, the ore is located 122 to 244 m from the contact between the limestone and adjacent gneisses. The ore body extends to a depth of 762 m.

The two ore bodies are presently separated by 5 km, but were from all indications at one time continuous.

The internal structure of the ore bodies is quite complex. It consists of lenslike or tabular layers of ore and scarn ranging from a few meters to 9.1 meters in thickness. Some of the thinner scarn bodies show boudinage structure. Small intrusive bodies of pegmatite may also be seen irregularly interspersed throughout the ore bodies. (See Map 1.)

The ore bodies display prominent faulting and jointing with the principle jointing direction being N 30°E. The joints often contain a variety of low-temperature hydrothermal minerals including Mn and Zn minerals as well as calcite and serpentine.

In general the ore is too thinly layered to be separately and selectively mined except at one unique location in the Sterling ore body where a cross vein cuts between the two limbs of the fold. Here a vein of high grade black willemite occurs.

The list of minerals found in the Franklin-Sterling Hill deposits follows in Table 2. The main ore consists of 40 percent franklinite, $(\text{Zn}, \text{Fe}, \text{Mn})\text{O} \cdot (\text{Fe}, \text{Mn})_2\text{O}_3$; 23 percent willemite, $2\text{ZnO} \cdot \text{SiO}_2$; and less than 1 percent zincite, ZnO . The remainder is a gangue of silicates and carbonates.

TABLE 2. Minerals of the Franklin and Sterling Hill Deposits

Actinolite: S	Hortonolite: O
Adamite: W	Hyalophane: S
Aegirine-augite: S	Hydrohetaerolite: W
Allactite: V	Hydrozincite: W
Alleghanyite: O, V	Johannsenite: S
Andradite: S, AS	Kaolinite: W
Anglesite: W	Kutnahorite: V
Anorthite: M	Lead: As, V
Anorthoclase: S	Leucophoenicite: O, V
Antigorite: AS, V	Linarite: W
Apatite: O, S	Loellingite: O
Aragonite: W	Magnesianiebeckite: V
Arsenopyrite: O, M	Magnetite: O
Augite: S	Malachite: W
Aurichalcite: W	Manganberzeliite: V
Axinite: AS, V	Manganosite: S
Azurite: W	Manganpyrosomalite: V
Bannisterite: S	Margarite: M
Barite: S, V	Margarosanite: AS
Barylite: S	McGovernite: AS, V
Barysilite: AS	Melanterite: W
Bementite: AS, V	Microcline: S
Biotite: S	Mooreite: AS, V
Bornite: V	Nasonite: AS, V
Brochantite: W	Nontronite: W
Brucite: V	Norbergite: M
Cahnite: V	Orthoclase: S
Calcite: O, S, AS, V, M	Paragasite: S
Celestite: V, AS	Pectolite: AS
Celsian: S	Phlogopite: M
Chalcocite: V	Prehnite: AS
Chalcophanite: W	Pyrite: V, M
Chalcopyrite: V	Pyrochroite: V
Chondrodite: M	Pyrrhotite: S, M
Chrysocolla: W	Quartz: S
Chrysotile: AS, V	Rhodochrosite: V
Clinohedrite: V	Rhodonite: O, S, V
Copper: V, AS	Roebingite: AS
Corundum: M	Roeppertite: O
Cryptomelane: W	Rosasite: W
Cumingtonite: S	Sauconite: W
Cuprite: V, W	Scapolite: M
Datolite: AS, V	Schallerite: V
Diopside: S, M	Scheelite: M
Djurleite: V	Serpentine: AS, V
Dolomite: V, M	Siderite: V
Edenite: S	Smithsonite: W
Epidote: S	Sonolite: O, S
Esperite: S	Spessartine: S
Ferroschallerite: S	Sphalerite: O, V
Fluoborite: V, AS	Spinel: M
Fluorite: O, S, V	Stilbite: V
Franklinite: O, S, AS, V	Stilpnomelane: V
Friedelite: S, V	Sussexite: S, V
Gahnite: S	Svabite: O, S
Galena: V	Talc: S, AS, M
Ganophyllite: AS, V	Tephroite: O
Glaucophane: S	Thomsonite: AS, V
Goethite: W	Titanite: M
Graphite: M	Todorokite: W
Grossular: S	Tourmaline (dravite): M
Hancockite: AS, V	Tremolite: S, M
Hardystonite: S	Vesuvianite: S
Hastingsite: S	Vredenburgite: O
Hedyphane: S, V	Willemite: O, S, AS, V
Hematite (specular): S, AS	Wollastonite: S
Hendricksite: S	Woodruffite: W
Hetaerolite: V	Xonotlite: AS
Hodgkinsonite: V	Zincite: O

Key to types of occurrence: O = primary ore, S = calcsilicate bodies, AS = hydrothermally altered calcsilicate bodies, V = hydrothermal veins, W = weathering products, M = Franklin marble.
The list excludes varietal terms and numerous rare minerals.

(After Frondel and Baum, p 170)

In the mineralogy, Fe^{+3} and Mn^{+2} predominate.

SUMMARY OF NEZILOVO

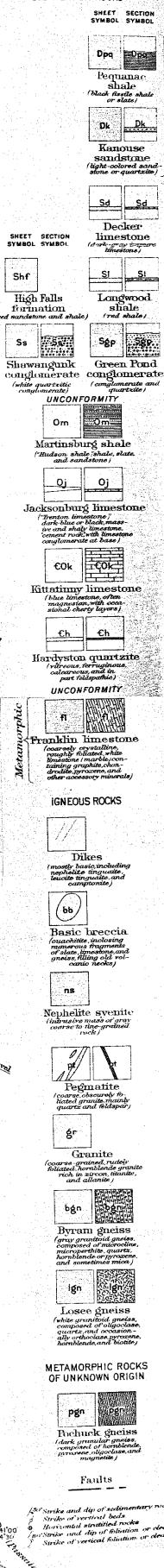
The village of Nezilovo lies in the southern political province of Yugoslavia known as the Peoples' Republic of Macedonia. This is 38 km southward of the Macedonian capital Skopje. (See Map 2.) A geologic formation exists here which is believed to be similar in origin and perhaps in mineralogy to the Franklin area. If similar deposits should be found in this area it would undoubtedly be a boon to Yugoslavia's available natural resources and economy. Little information was available concerning this region. Different sources in Yugoslavia were asked for additional information, but as yet none have responded. From the general structure map of Yugoslavia the deposit would seem to be located in an anticlinal fold plunging north-northwestward. (See Map 3.) The area seems to be intensely folded, faulted and metamorphosed. This would certainly seem to be a likely candidate for mineralization. It is presumed that no mining is currently taking place in the area.

METHODS OF APPROACH

For this research, six hand specimens were obtained from Dr. Ernest Ehlers who collected these specimens on a trip to Yugoslavia. The specimens were collected within 12 m of each other.

Small portions of each rock were crushed and

SEDIMENTARY ROCKS



Geology by Harry B. Kimmel,
Arthur C. Spencer, and Stuart Weller.
Surveyed in 1892-1897 and 1905.

- 1/2° Strike and dip of sedimentary rocks
- ° Strike of vertical beds
- 0 Horizontal stratified rocks
- 1/2° Strike and dip of foliation or cleavage
- ° Strike of vertical foliation or cleavage

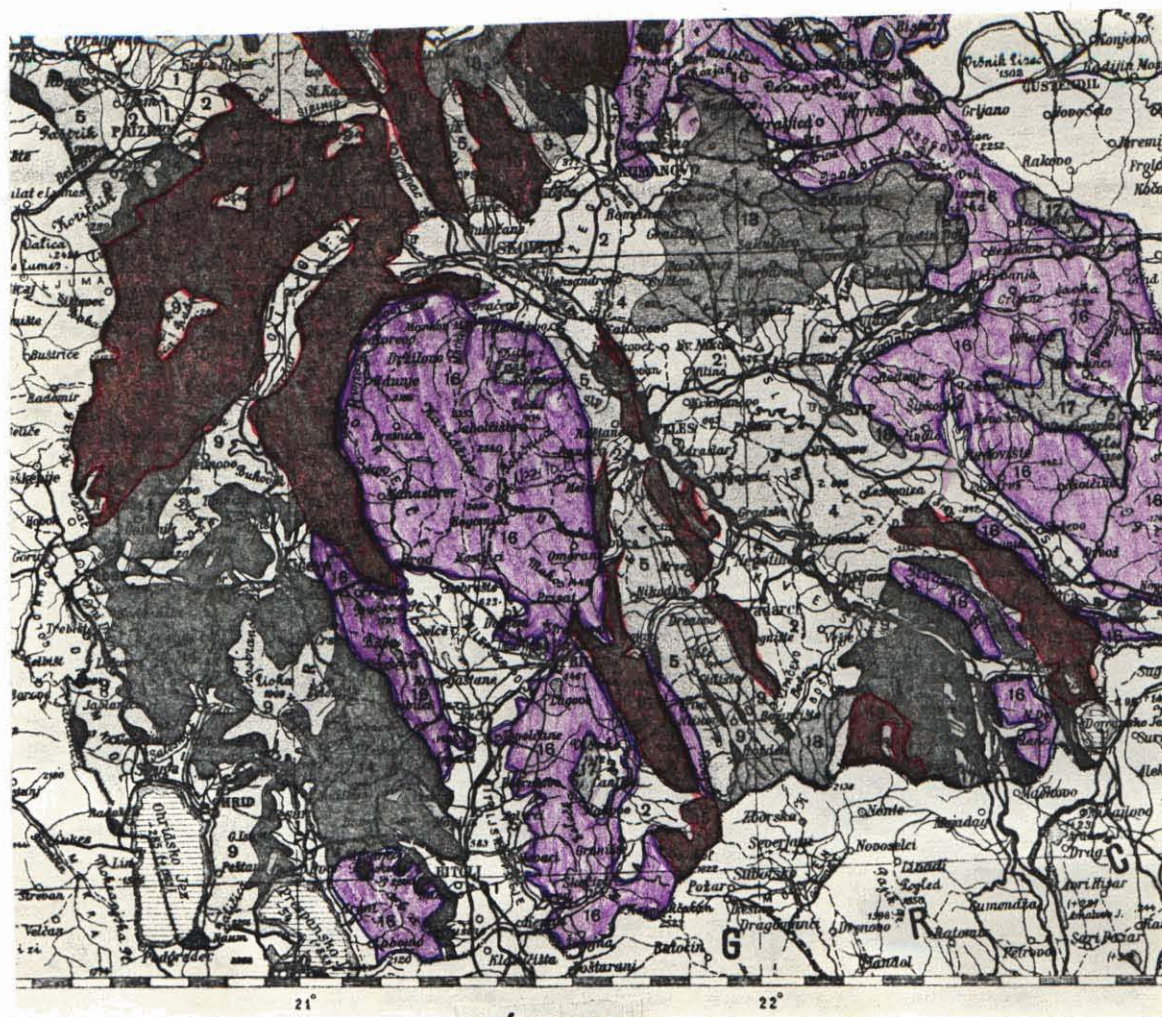


- Автомобилски пат I ред
- Автомобилски пат II ред
- Автомобилски пат III ред

КОНУСНА ПРОЈЕКЦИЈА
1:500,000

Км 0 10 20 30 40 50

Издание и издавачка куќа БЕСКАРТА



LÉGENDE.

1	Quaternaire.	11	Phénites.
2	Néogène.	12	Permien (grès rouges).
3	Les couches oligocènes d'eaux douces.	13	Permo-carbonifère.
4	Paléogène.	14	Carbonifère.
5	Crétacé supérieur.	15	Paléozoïque en général (schistes cristallins du II ^{me} groupe).
6	Crétacé inférieur.	16	Schistes cristallins (du I ^{er} groupe).
7	Jurassique moyen et supérieur.	17	Granites.
8	Jurassique inférieur. (Lias).	18	Andézites et roches trachytiques.
9	Trias moyen et supérieur.	19	Diabases et melaphyres.
10	Trias inférieur.	20	Gabbros et serpentines.

IZDANJE KNJIZARE FRANJE BACHA, BEOGRAD (1930-1931)
Geološka Karta Kraljevine Jugoslavije

MAP 3

separated. Nonopaque minerals were then crushed to between 100 and 200 mesh size and analysed optically. Opaque and semi-opaque minerals were crushed to less than 325 mesh and were analysed by X-ray powder diffraction method. (See Table 3.) The X-ray machine was equipped with a copper tube with nickel shielding. The machine was set on a coarse adjustment of 1500 volts in the high voltage selector and on 35 kilovolts and 15 milliamps for the power generation. The crushed material was moistened with a water and Duco Cement solution to afix it to the slide for analysis.

Baric (1960) reports the presence and characteristics of gahnite, piedmontite, and rutile from this locale.

In this investigation the following minerals were isolated: albite, $\text{Na}_2\text{O} \cdot \text{Al}_2\text{O}_3 \cdot 6\text{SiO}_2$; alurgite, $\text{K}_2\text{MgAl}_4\text{Si}_7\text{O}_{22}(\text{OH})_2$; barite, BaSO_4 ; biotite, $\text{K}_2(\text{Fe}, \text{Mn}, \text{Mg})_4(\text{Al}, \text{Fe})_4\text{Si}_6\text{O}_{22}\text{OH}_2$; dolomite, $\text{CaO} \cdot \text{MgO} \cdot 2\text{CO}_2$; epidote, $4\text{CaO} \cdot 3(\text{Al}, \text{Fe})_2\text{O}_3$; gahnite, $(\text{Zn}, \text{Mn})\text{O} \cdot \text{Al}_2\text{O}_3$; hematite, Fe_2O_3 ; magnetite, Fe_3O_4 ; microcline, $\text{K}_2\text{O} \cdot \text{Al}_2\text{O}_3 \cdot 6\text{SiO}_2$; muscovite, $(\text{Na}, \text{K})_2\text{Al}_6\text{Si}_6\text{O}_{22}(\text{OH})_4$; piedmontite, $4\text{CaO} \cdot 3(\text{Al}, \text{Fe}, \text{Mn})_2\text{O}_3 \cdot 6\text{SiO}_2 \cdot \text{H}_2\text{O}$; quartz, SiO_2 ; rutile, TiO_2 ; and talc, $3\text{MgO} \cdot 4\text{SiO}_2 \cdot \text{H}_2\text{O}$. (See Table 4.)

CONCLUSIONS

All of the minerals found in this investigation from the Nezilovo area have been found at Franklin with the exception of piedmontite and alurgite. More attention should be paid to the minerals that were not present, particularly the ore minerals which occur at Franklin: franklinite, willemite, and zincite.

RESULTS OF X-RAY ANALYSIS

Mineral	Composition	DEGREES ROTATION INSPECTED	PEAK LOCATION	λ , α	RELATIVE INTENSITY
GANNITE	$(Zn, Mn)O \cdot Al_2O_3$	$5^\circ - 82^\circ$	36.7°	2.4487	100%
			59.0°	1.5655	85%
			31.2°	2.8666	65%
			64.9°	1.4367	35%
			65.0°	1.4347	30%
			55.4°	1.6584	20%
			77.0°	1.2383	15%
			44.6°	2.0316	10%
			77.2°	1.2356	10%
			44.8°	2.0230	8%
HEMATITE \bar{c} MAGNETITE	Fe_2O_3 Fe_3O_4	$5^\circ - 60^\circ$	33.1°	2.7063	100%
			49.4°	1.8448	60%
			54.0°	1.6980	55%
			35.7°	2.5149	50%
			24.1°	3.6926	40%
			27.9°	3.1977	35%
			46.5°	1.9529	25%
			40.8°	2.2116	25%
DOLOMITE WITH IRON	$CaO \cdot MgO \cdot 2CO_2$	$25^\circ - 76^\circ$	30.9°	2.8938	100%
			51.0°	1.7906	40%
			50.5°	1.8072	25%
			41.1°	2.2012	22%
			35.3°	2.5425	10%
			33.5°	2.6749	10%
			44.9°	2.0187	9%
			65.1°	1.4328	8%
			63.4°	1.4670	8%
			59.8°	1.5465	5%

TABLE 3.

NEŽILOVO MINERALS

Mineral	X-RAY ANALYSIS DONE	Composition	β or ω Index	2 γ	SIGN	Biref.	Cleavage	Microscopic Color	Microscopic Color	misc.
1. Albite		$\text{Na}_2\text{O} \cdot \text{Al}_2\text{O}_3 \cdot 6\text{SiO}_2$	1.53 β	74°	+	.011	{100} {100} PERF.	WHITE	WHITE-CLEAR	POSSIBLY TWIN. SOME GRAINS
2. Analcite		$\text{K}_2\text{MgAl}_4\text{Si}_7\text{O}_{22}(\text{OH})_2$	1.59 β	~15°	—	.046	{100} {100} mic	RED	PINK	
3. Barite		BaSO_4	1.64 β	37°	+	.013	{100} {110} {100} PERF. {100} IMP.	WHITE	WHITE-CLEAR	
4. Biotite		$\text{K}_2(\text{Fe}, \text{Mn}, \text{Mg})_2(\text{Al}, \text{Fe})_2\text{Si}_6\text{O}_{20}$	1.63 β	~10°	—	.050	{100} {100} {100} {100} {100} {100}	BLACK	ALMOST OPAQUE BLACK	
5. Dolomite	*	$\text{CaO} \cdot \text{MgO} \cdot \text{CaCO}_3$	1.68 ω	0°	—	.181	{110} {110} PERF.	BROWN	CLEAR WITH BROWN STAIN	IN SOME CASES MIN. LIPS CLEAR WITH PINK CRUST
6. Epidote		$4\text{CaO} \cdot 3(\text{Al}, \text{Fe})_2\text{O}_3$	1.74 β	80°	—	.038	{100} {100} PERF. {100} IMP.	DK. GREEN	LT. GREEN	
7. Garnet		$(\text{Zn}, \text{Mn})\text{O} \cdot \text{Al}_2\text{O}_3$	1.78	—	ISOTROPIC	—	None	Green, Blue, Black	Green, Blue, Brown, Green	VARIED COMPOSITION
8. Hematite		Fe_2O_3	3.01 ω	0°	—	.28	{100} {100} PERF.	BLACK	Bt. RED	ALTERATION FROM MAGNETITE
9. Magnetite	*	Fe_3O_4	—	—	OPAQUE	—	—	BLACK	BLACK	Magnetite
10. Microcline		$\text{K}_2\text{O} \cdot \text{Al}_2\text{O}_3 \cdot 6\text{SiO}_2$	1.53 β	83°	—	.004	{100} {100} PERF.	LT. BLUE	WHITE-CLEAR	TACTILE TWINNING
11. Muscovite		$(\text{Na}, \text{K})_2\text{Al}_2\text{Si}_6\text{O}_{20}(\text{OH})_4$	1.59 β	42°	—	.043	{100} {100} mic	YELLOWISH	Clear to yellow green	
12. Pyromorphite	*	$4\text{CaO} \cdot 3(\text{Al}, \text{Fe}, \text{Mn})_2\text{O}_3 \cdot 6\text{SiO}_2 \cdot \text{H}_2\text{O}$	1.77 β	68°	+	.018	{100} {100} PERF. {100} {100} {100} {100}	BLACK	Green-Brown	Epidote Group
13. Quartz		SiO_2	1.54 ω	0°	+	.009	None	Clear	Clear	
14. Rutile		TiO_2	2.62 ω	0°	+	.287	{100} {110} {100} DISTINCT	BLACK	DK. BROWN	
15. Talc		$3\text{MgO} \cdot 4\text{SiO}_2 \cdot \text{H}_2\text{O}$	1.59 β	~15°	—	.050	{100} {100} mic	WHITE	WHITE-CLEAR	FOULATIONS IN GRAINS

TABLE 4

Baric (1960) lists Nezilova as a lead-zinc deposit. No lead minerals and only one zinc mineral, gahnite, were found in this investigation. Due to the limited amount of work previously done in the area and the unavailability at this time of more rock specimens to be analysed, it is impossible to tell how closely these deposits are related, though such a relationship might certainly be possible. A need for further investigation is indicated and might prove quite fruitful.

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